

SOME SOLUTIONS TO EQUATIONS OF  
MOTION IN THE EQUATORIAL  
REGIONS

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R. M. JONSON

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IN THE EQUATORIAL REGIONS

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R. M. Jonson





SOME SOLUTIONS TO EQUATIONS OF MOTION  
IN THE EQUATORIAL REGIONS

by  
Russell Martin Jonson  
Lieutenant, United States Navy

Submitted in partial fulfillment  
of the requirements  
for the degree of  
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IN AEROSPACE

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Monterey, California  
1952



This work is accepted as fulfilling  
the thesis requirements for the degree of

MASTER OF SCIENCE  
IN AEROLOGY

from the  
United States Naval Postgraduate School



## PREFACE

This investigation was conducted at the United States Naval Postgraduate School, Monterey, California as the thesis requirements for the degree of Master of Science in Aerology.

For help and advice received in its preparation the author is indebted to Associate Professor G. J. Haltiner of the U. S. Naval Postgraduate School.



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Summary of the

10

1. The first part of the report is devoted to a general survey of the situation in the country.	10
2. The second part is devoted to a detailed examination of the various branches of the economy.	20
3. The third part is devoted to a study of the social and cultural conditions of the population.	30
4. The fourth part is devoted to a study of the political and administrative organization of the country.	40
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7. The seventh part is devoted to a study of the scientific and literary achievements of the country.	70
8. The eighth part is devoted to a study of the art and architecture of the country.	80
9. The ninth part is devoted to a study of the history and traditions of the country.	90
10. The tenth part is devoted to a study of the future prospects of the country.	100

# TABLE OF SYMBOLS AND ABBREVIATIONS

$\lambda$	Coriolis parameter = $2\omega \sin \phi$
$\rho$	Density of air
$t$	Time
$\phi$	Latitude
$\omega$	Angular speed of the earth = $7.29 \times 10^{-5} \text{ sec}^{-1}$
$x$	Coordinate axis - East
$y$	Coordinate axis - North
$u$	Easterly velocity component
$v$	Northerly velocity component
$a$	Radius of the earth = 6378 km
$T$	Period of pressure variation = 12 hours
$C_1, C_2$	Constant of integration



## I. INTRODUCTION

In mid-latitudes the geostrophic and gradient solutions of the equations of motion generally give a good approximation to the true wind. However, in the equatorial regions the acceleration terms, both centripetal and tangential, are frequently of the same order of magnitude as the pressure and Coriolis forces; hence the gradient and geostrophic solutions are not generally applicable.

To render the equations integrable various simplifications have been used. Notable among these are the assumptions that the flow is (a) horizontal, (b) frictionless, (c) homogeneous, (d) non-divergent and (e) steady. Some investigators have neglected the Coriolis force, while others have either treated it as a constant or replaced  $\sin \phi$  by  $\phi$ . Practically all authors assume (a) through (c) above as will be done by all discussions in this paper.

In addition to the above assumptions, Grimes [5] let motion be independent of longitude. He allowed for variation of the Coriolis force by assuming  $\sin \phi = \phi$  with  $\phi$  as a function of  $y$ . His solutions gave particular streamlines which were determined from initial values of velocity and vorticity. Pressure distributions were also determined from these solutions; however, it is generally preferable to derive motions from known pressure distributions and initial velocities. Grimes obtained isobaric patterns which were similar to mean maps in the Indian and South Pacific oceans but pressure gradients were less



than those usually observed. Crossley [2] found this solution to apply only to mean motions and not to the instantaneous flow pattern.

Crossley [1] removed the restriction on independence of longitude and corrected the fault that all isobars crossed the equator at right angles in the above solution. The same author [3] obtained a solution in spherical coordinates.

Grimes' original assumptions less the requirement of non-divergence were used in Chapter I of this paper.

Forsdyke [4] took a constant Coriolis force with motion at any instant everywhere the same, that is, the space derivatives do not appear in acceleration terms. He assumed a pressure distribution of the form  $p = p_0 + A(t)y$  for various functions  $A(t)$ .

By expressing the resulting motion as geostrophic plus ageostrophic components, he suggests a method for relating wind to the pressure distribution in the tropics.

In Chapter II an attempt is made to generalize Forsdyke's solution by making less restrictions on the Coriolis parameter and assuming a pressure distribution of the form  $p = p_0 + x \sin \frac{2\pi t}{T}$ .

In Chapter III a pressure field of the form  $p = p_0 + Ae^{-\epsilon t} B(1 - e^{-\gamma t})y$  where  $A$ ,  $B$ ,  $\epsilon$  and  $\gamma$  are constants is discussed. This pressure field was suggested but not developed by Forsdyke [4].

More recently, Schmidt [7] concentrated on the kinematics of flow by starting with Rossby's vorticity equation and assuming the dependent variables to be functions of  $y$  only. He obtained values for  $u$  in terms







of  $y$  assuming  $v$  known and plotted streamlines for a monsoon current, high pressure at the equator (doldrums), and a line of convergence stretching in an east-west direction along the equator. Some of the streamlines obtained in Chapter II of this paper closely resemble Schmidt's.

Grimes [6] using his original assumptions less independence with longitude, obtains a set of pseudo-geostrophic equations where streamlines are absolute vorticity trajectories. They are parallel to isobars of "dynamic pressure" ( $P = p + \frac{1}{2} \rho V^2$ , where  $V^2 = u^2 + v^2$ ). The difficulty in this approach is the necessity for drawing accurate isobars of dynamic pressure, which must be drawn at intervals of one-fifth of a millibar. This is particularly difficult where reduction to sea level is necessary.

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## II. MOTION AS A FUNCTION OF $y$ ALONE

The equations of horizontal frictionless motion are:

$$\frac{du}{dt} - 2\omega v \sin\phi = -\frac{1}{\rho} \frac{\partial p}{\partial x} \quad (1)$$

$$\frac{dv}{dt} + 2\omega u \sin\phi = -\frac{1}{\rho} \frac{\partial p}{\partial y} \quad (2)$$

In this solution steady-state is assumed with motion independent of longitude ( $\frac{\partial u}{\partial x} = \frac{\partial v}{\partial x} = 0$ ). Pressure is a linear function of latitude ( $y$ ) alone with isobars equally spaced and parallel to the equator. The Coriolis force is treated in a similar manner to Grimes [5] by letting  $\sin\phi = \phi = y/a$ . Thus the Coriolis parameter is  $Ky$  where  $K = 2\omega/a$ .

With these assumptions the equations of motion become:

$$v \frac{\partial u}{\partial y} = Kyv \quad (3)$$

$$v \frac{\partial v}{\partial y} = K \frac{\partial p}{\partial y} - Kyu \quad (4)$$

Equation (3) readily integrates to

$$u = \frac{1}{2} Ky^2 + u_0 \quad (5)$$

Substitution of this value of  $u$  into equation (4) and integrating yields the following expression for  $v$ ,

Find the area of the region bounded by the curves

$y = \sin x$  and  $y = \cos x$  from  $x = 0$  to  $x = \frac{\pi}{2}$ .

$$(1) \quad \frac{1}{2} \left( \frac{\pi}{2} - 0 \right) = \frac{\pi}{4} \approx 0.7854$$

$$(2) \quad \frac{1}{2} \left( \frac{\pi}{2} - 0 \right) = \frac{\pi}{4} \approx 0.7854$$

Therefore, the area of the region bounded by the curves  $y = \sin x$  and  $y = \cos x$  from  $x = 0$  to  $x = \frac{\pi}{2}$  is

$\frac{\pi}{4} \approx 0.7854$ . (Note: The area is the same for both regions.)

Example 2: Find the area of the region bounded by the curves  $y = \sin x$  and  $y = \cos x$  from  $x = \frac{\pi}{2}$  to  $x = \pi$ .

Solution: The curves  $y = \sin x$  and  $y = \cos x$  intersect at  $x = \frac{\pi}{2}$  and  $x = \frac{3\pi}{2}$ . The region bounded by the curves from  $x = \frac{\pi}{2}$  to  $x = \pi$  is shown in Figure 6.10.

Since  $y = \sin x$  is above  $y = \cos x$  on the interval  $[\frac{\pi}{2}, \pi]$ , the area of the region is given by

$$A = \int_{\frac{\pi}{2}}^{\pi} (\sin x - \cos x) dx$$

Evaluating the integral, we get

$$(1) \quad A = \left[ -\cos x - \sin x \right]_{\frac{\pi}{2}}^{\pi} = (-1 - 0) - (0 - 1) = 0$$

$$(2) \quad A = \left[ -\cos x - \sin x \right]_{\frac{\pi}{2}}^{\pi} = (-1 - 0) - (0 - 1) = 0$$

Therefore, the area of the region bounded by the curves  $y = \sin x$  and  $y = \cos x$  from  $x = \frac{\pi}{2}$  to  $x = \pi$  is 0.

$$(3) \quad A = \left[ -\cos x - \sin x \right]_{\frac{\pi}{2}}^{\pi} = (-1 - 0) - (0 - 1) = 0$$

Therefore, the area of the region bounded by the curves  $y = \sin x$  and  $y = \cos x$  from  $x = \frac{\pi}{2}$  to  $x = \pi$  is 0.

Example 3: Find the area of the region bounded by the curves  $y = \sin x$  and  $y = \cos x$  from  $x = 0$  to  $x = \frac{3\pi}{2}$ .

$$v^2 = v_0^2 + 2 \alpha (p - p_0) - \frac{1}{4} K^2 y^4 - K u_0 y^2 \quad (6)$$

$u_0, v_0, p_0$  being values of velocity and pressure at  $y = 0$ .

Numerical values of the slope of the streamlines were found at every 500 kilometers with varying initial conditions and a constant pressure gradient of 1.5 millibars per 500 km and a specific volume of 860 cm<sup>3</sup>/gm.

These streamline patterns (see Figures 1-5) resemble observed mean motions in the Indian and Southern Pacific oceans. Values of  $v$  for  $y > 1500$  km were imaginary indicating steady-state is no longer valid at that distance from the equator with the given initial conditions.

Most of the faults of Grimes' and Crossley's solutions still remain in this simple pattern, the most important being that steady state is not generally applicable and therefore solutions apply only to mean motions. It should also be noted that the determination of the streamline patterns from even these simple pressure distributions is a laborious process.

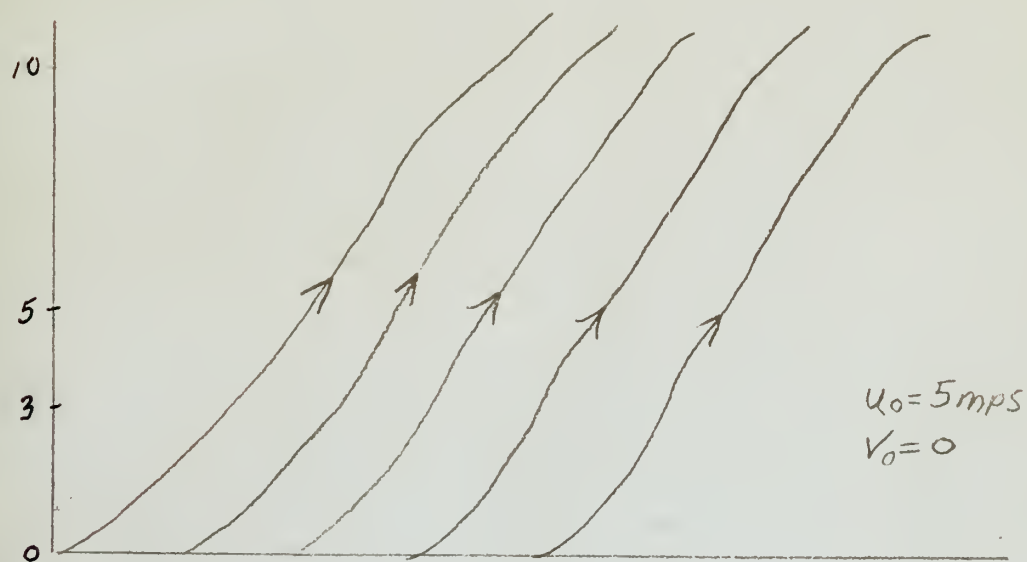
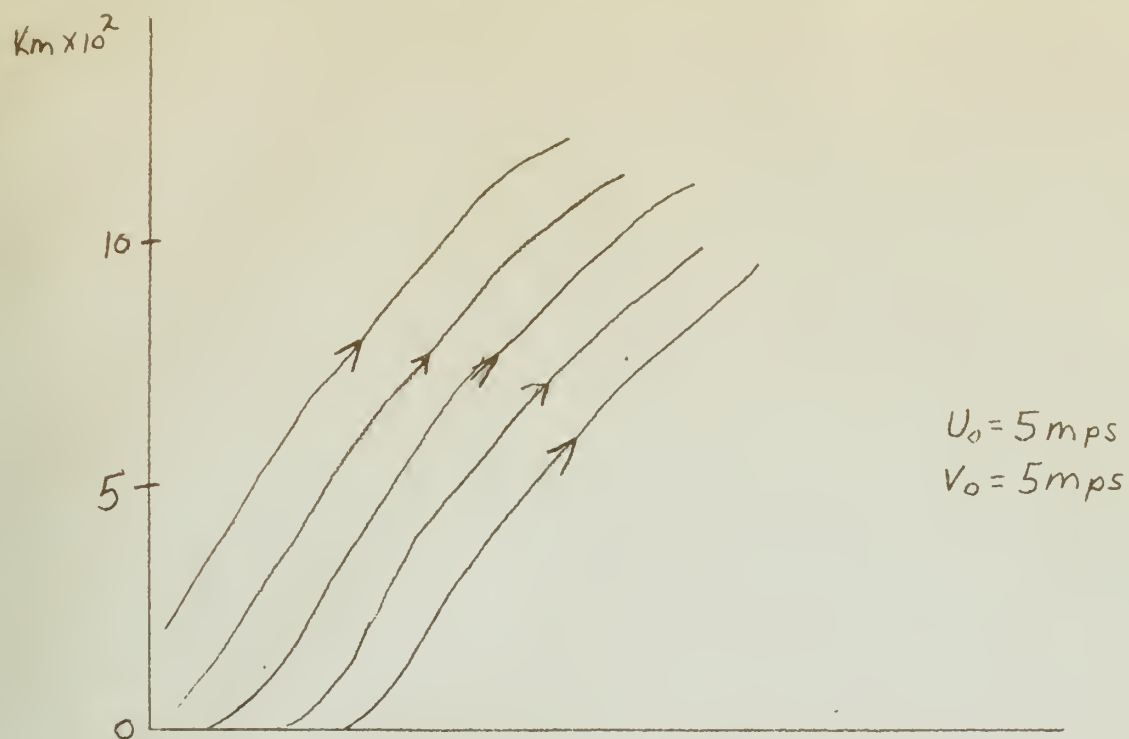
However, one improvement is the removal of the restriction of no horizontal divergence. Zones of divergence and convergence are known by observations to exist in the equatorial regions.

The first of these is the fact that the United States is a young nation, and its history is therefore a history of growth and development. The second is the fact that the United States is a large nation, and its history is therefore a history of expansion and conquest. The third is the fact that the United States is a diverse nation, and its history is therefore a history of conflict and compromise.

The fourth is the fact that the United States is a nation of immigrants, and its history is therefore a history of assimilation and adaptation. The fifth is the fact that the United States is a nation of pioneers, and its history is therefore a history of exploration and discovery. The sixth is the fact that the United States is a nation of entrepreneurs, and its history is therefore a history of innovation and progress.

The seventh is the fact that the United States is a nation of idealists, and its history is therefore a history of aspiration and achievement. The eighth is the fact that the United States is a nation of pragmatists, and its history is therefore a history of compromise and adaptation. The ninth is the fact that the United States is a nation of optimists, and its history is therefore a history of hope and progress. The tenth is the fact that the United States is a nation of pessimists, and its history is therefore a history of despair and decline.

The eleventh is the fact that the United States is a nation of dreamers, and its history is therefore a history of vision and ambition. The twelfth is the fact that the United States is a nation of doers, and its history is therefore a history of action and achievement. The thirteenth is the fact that the United States is a nation of thinkers, and its history is therefore a history of reflection and wisdom. The fourteenth is the fact that the United States is a nation of feelers, and its history is therefore a history of emotion and passion.



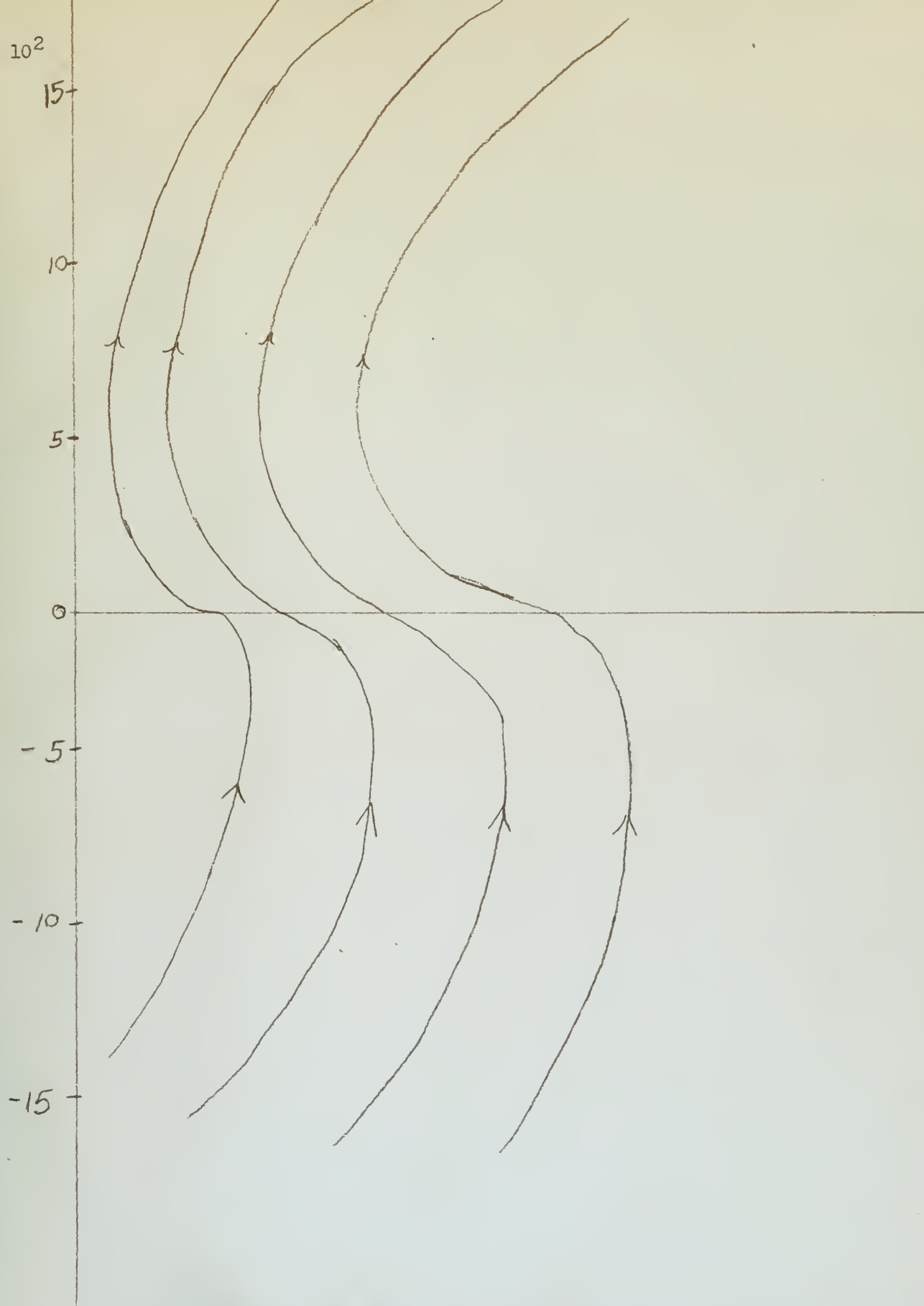
Sea level streamlines for  $u_0 = 5 \text{ mps}$ ,  
 $v_0 = 5 \text{ mps}$  and for  $u_0 = 5 \text{ mps}$ ,  $v_0 = 0$  at  $y = 0$

Fig. 1





km  $\times 10^2$

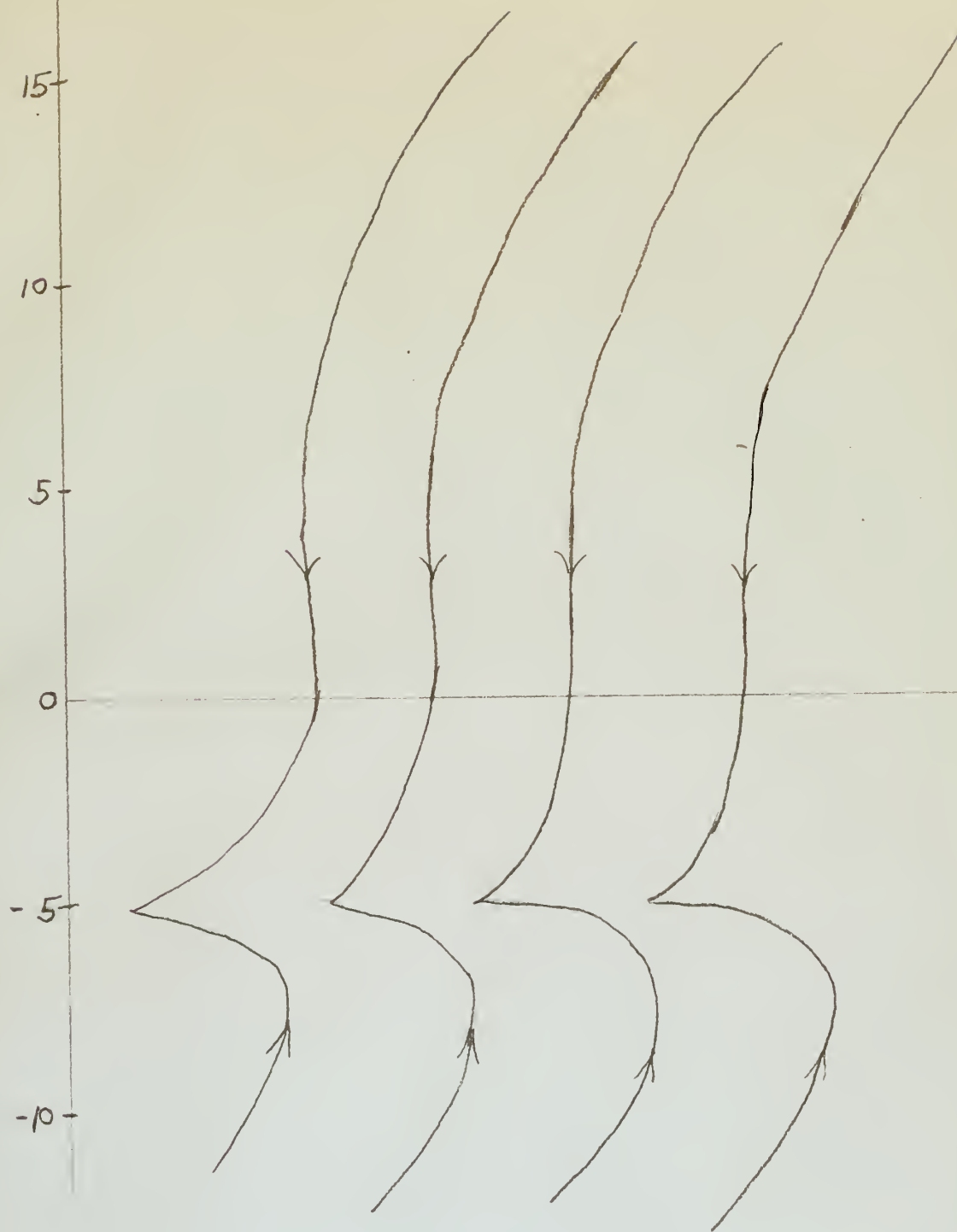


Sea level streamlines for  
 $u_0 = -5$  mps,  $v_0 = 0$  at  $y = 0$

Fig. 2



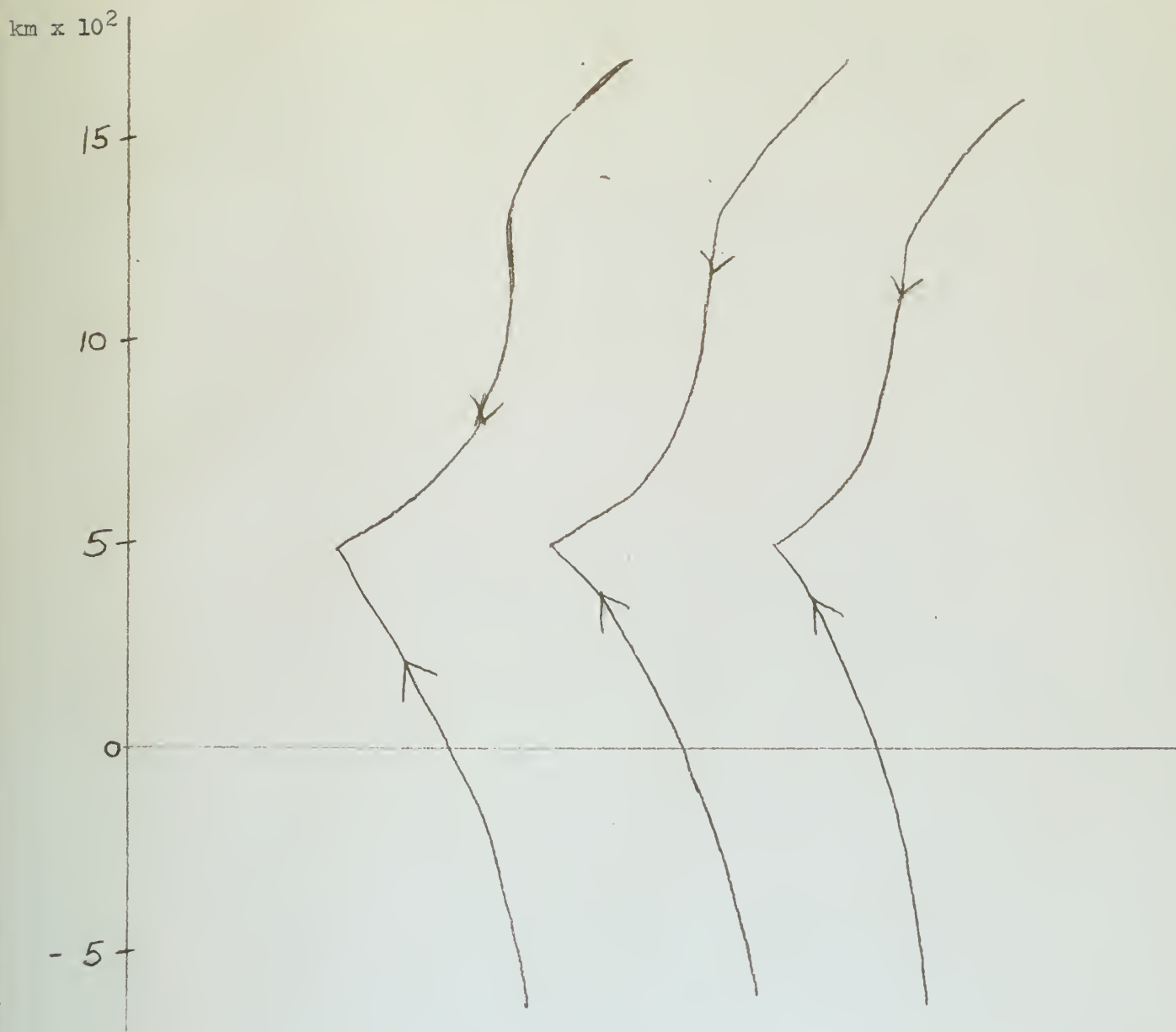
km x 10<sup>2</sup>



Sea level streamlines for  
 $u_0 = -5$  mps,  $v_0 = 0$  at  $y = -500$  km

Fig. 3

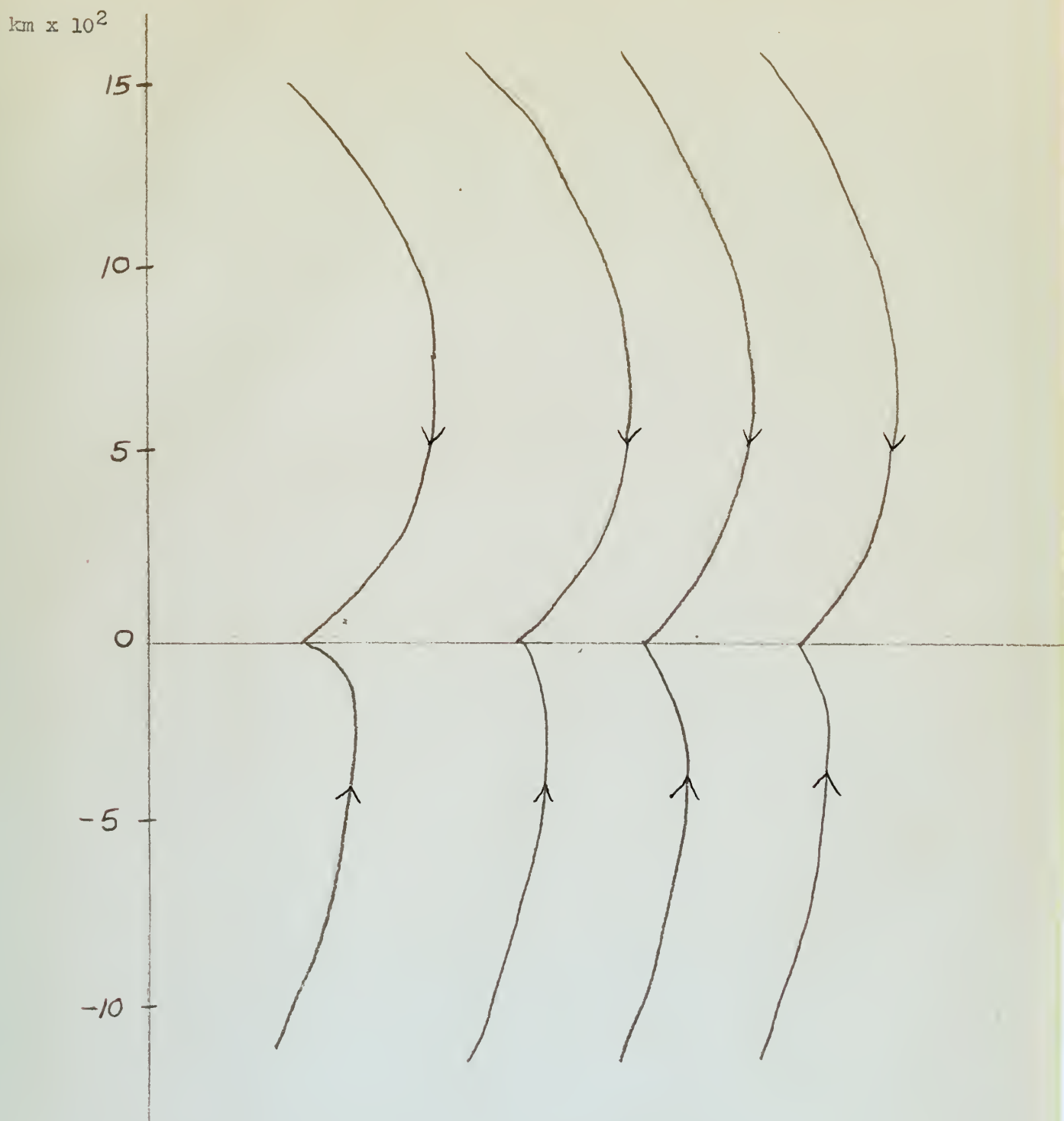




Sea level streamlines for  
 $u_0 = -5$  mps,  $v_0 = -2$  mps at  $y = 500$  km

Fig. 4





Sea level streamlines for  
 $u_0 = -5$  mps,  $v_0 = -5$  mps at  $y = 0$

Fig. 5





### III. SPATIALLY INDEPENDENT MOTION WITH VARIABLE CORIOLIS PARAMETER

In this approach Forsdyke's [4] assumptions were used but the Coriolis parameter was varied to secure a more general solution. Motions were considered uniform which is probably valid over areas five degrees square. The pressure distribution  $p = p_0 + p_1 X \sin \frac{2\pi t}{T}$  was assumed.

The equations of motion with the above pressure distribution are:

$$\frac{du}{dt} = \lambda v - \frac{1}{\rho} \frac{\partial p}{\partial x} = \lambda v - \frac{p_1}{\rho} \sin \frac{2\pi t}{T} \quad (7)$$

$$\frac{dv}{dt} = -\lambda u \quad (8)$$

Differentiating (8) with respect to  $t$ , we obtain

$$\frac{d^2v}{dt^2} = -\lambda \frac{du}{dt} - u \frac{d\lambda}{dt} \quad (9)$$

moreover  $\lambda = 2\omega \sin \phi$  ;  $\frac{d\lambda}{dt} = 2\omega \cos \phi \frac{d\phi}{dt}$  ;  $\phi = y/a$ ; and

$\frac{d\phi}{dt} = \frac{1}{a} dy/dt = v/a$ . Since  $\cos \phi$  varies only 20% from  $0^\circ$  to  $20^\circ$ , we assume  $\cos \phi = 1$  and it follows that,

$$\frac{d\lambda}{dt} = \frac{2\omega v}{a} \quad (10)$$

Substituting the expressions for  $du/dt$ ,  $u$ , and  $d\lambda/dt$  from equations (7), (8), and (10) respectively, we get



$$\frac{d^2 v}{dt^2} - \frac{2\omega v}{a\tau} \frac{dv}{dt} + \lambda^2 v = \frac{\lambda P}{e} \sin \frac{2\pi t}{T} \quad (11)$$

To solve this differential equation a mean value of  $v(\bar{v})$  was assumed giving a second order equation with constant coefficients.

The solution with constants of integration evaluated for the boundary condition,  $v = 0$  at  $t = 0$  and  $C_1 = 1$  is

$$\begin{aligned} v = & e^{\frac{\omega \bar{v} + \sqrt{\omega^2 \bar{v}^2 - a^2 \lambda^4}}{a\lambda} t} + (-1-B) e^{\frac{\omega \bar{v} - \sqrt{\omega^2 \bar{v}^2 - a^2 \lambda^4}}{a\lambda} t} \\ & + A \sin \frac{2\pi t}{T} + B \cos \frac{2\pi t}{T} \\ u = & - \left( \frac{\omega \bar{v} + \sqrt{\omega^2 \bar{v}^2 - a^2 \lambda^4}}{a\lambda^2} \right) e^{\frac{\omega \bar{v} - \sqrt{\omega^2 \bar{v}^2 - a^2 \lambda^4}}{a\lambda} t} \\ & + (1+B) \left( \frac{\omega \bar{v} - \sqrt{\omega^2 \bar{v}^2 - a^2 \lambda^4}}{a\lambda^2} \right) e^{\frac{\omega \bar{v} + \sqrt{\omega^2 \bar{v}^2 - a^2 \lambda^4}}{a\lambda} t} \\ & - \frac{A}{\lambda} \frac{2\pi}{T} \cos \frac{2\pi t}{T} + \frac{B}{\lambda} \frac{2\pi}{T} \sin \frac{2\pi t}{T} \end{aligned}$$

where

$$A = \frac{\lambda P (\lambda^2 T^2 - 4\pi^2)}{e T^2 \left[ \left( \frac{\lambda^2 T^2 - 4\pi^2}{T^2} \right)^2 + \left( \frac{4\pi \omega \bar{v}}{a\lambda T} \right)^2 \right]}$$

and

$$B = \frac{P, 4\pi \omega \bar{v}}{e a T \left[ \left( \frac{\lambda^2 T^2 - 4\pi^2}{T^2} \right)^2 + \left( \frac{4\pi \omega \bar{v}}{a\lambda T} \right)^2 \right]}$$

The special case for which  $\bar{v} = 0$  gives Forsdyke's solution. Assigning normal numerical values led to imaginary values of  $u$  and  $v$  and hence the analysis was carried no further. Some numerical constants will lead to real values for  $u$  and  $v$ ; however, the results do not conform to observed patterns.

$$x = \frac{1}{2} + \frac{1}{2}i + \frac{1}{2}j + \frac{1}{2}k$$

Let  $x = \frac{1}{2} + \frac{1}{2}i + \frac{1}{2}j + \frac{1}{2}k$  be a solution of the system of equations  $x^2 + y^2 + z^2 = 1$  and  $x + y + z = 1$ . Then  $x, y, z$  are the roots of the equation  $t^3 - 1 = 0$ .

$$x^2 + y^2 + z^2 = 1 \quad (1)$$

$$x + y + z = 1 \quad (2)$$

$$x^2 + y^2 + z^2 - (x + y + z)^2 = 1 - 1^2 = 0$$

$$x^2 + y^2 + z^2 - x^2 - y^2 - z^2 - 2xy - 2yz - 2zx = 0$$

$$-2xy - 2yz - 2zx = 0$$

$$xy + yz + zx = 0$$

$$\frac{x^2 + y^2 + z^2}{x + y + z} = \frac{1}{1} = 1$$

$$\frac{x^2 + y^2 + z^2}{x + y + z} = \frac{1}{1} = 1$$

Let  $x = \frac{1}{2} + \frac{1}{2}i + \frac{1}{2}j + \frac{1}{2}k$  be a solution of the system of equations  $x^2 + y^2 + z^2 = 1$  and  $x + y + z = 1$ . Then  $x, y, z$  are the roots of the equation  $t^3 - 1 = 0$ . The roots of this equation are  $1, \omega, \omega^2$ , where  $\omega = \frac{-1 + \sqrt{-3}}{2}$  and  $\omega^2 = \frac{-1 - \sqrt{-3}}{2}$ . Therefore, the solutions of the system are  $(1, \omega, \omega^2)$  and  $(1, \omega^2, \omega)$ .

#### IV. SPATIALLY INDEPENDENT MOTION FOR A CERTAIN PRESSURE DISTRIBUTION

Forsdyke [4] suggested a pressure distribution of the form  $p = p_0 + Ae^{-\epsilon t} x + B(1-e^{-\gamma t})y$  - where  $A, B, \epsilon, \gamma$  are constants to be applied to the spatially independent constant Coriolis parameter solution. He states this solution might be applied to curved isobars.

This statement appears to be inconsistent for at any time  $t$ , the pressure is a linear function of  $x$  and  $y$  giving an isobaric pattern consisting of straight parallel lines with their orientation changing with time. At  $t = 0$  the isobars are a function of  $x$  and perpendicular to the equator. As  $t$  increases and approaches infinity the isobars rotate and become parallel to the equator and are functions of  $y$  alone.

The constants  $A$  and  $B$  are the pressure gradients at times  $t = 0$  and  $t = \infty$  respectively. The constants  $\epsilon$  and  $\gamma$  determine how rapidly the pressure gradients change in the  $x$  and  $y$  directions.

In applying this pressure distribution  $\epsilon$  and  $\gamma$  were made equal for simplification.

The equations of motion with Forsdyke's assumption and the above pressure distribution become

$$\frac{du}{dt} = \lambda v - \frac{A}{\rho} e^{-\gamma t} \quad (12)$$

$$\frac{dv}{dt} = -\lambda u - \frac{B}{\rho}(1-e^{-\gamma t}) \quad (13)$$



The general solution of this set of differential equations consists of the complementary functions

$$u = C_2 \sin \lambda t - C_1 \cos \lambda t$$

$$v = C_1 \sin \lambda t + C_2 \cos \lambda t$$

and the particular integrals

$$u = \frac{B}{\rho \lambda} - \frac{(B\lambda + A\sigma)}{\rho(\sigma^2 + \lambda^2)} e^{-\sigma t}$$

$$v = \frac{A\lambda - B\sigma}{\rho(\sigma^2 + \lambda^2)} e^{-\sigma t}$$

If  $C_1$  and  $C_2$  are evaluated at the boundary conditions  $u = u_0$ ,  $v = v_0$  at  $t = 0$  and the values of geostrophic wind

$$u_g = -\frac{1}{\rho \lambda} \frac{\partial p}{\partial y} = -\frac{B}{\rho \lambda} (1 - e^{-\sigma t})$$

$$v_g = \frac{1}{\rho \lambda} \frac{\partial p}{\partial x} = \frac{A}{\rho \lambda} e^{-\sigma t}$$

are inserted we have

$$u = \left[ \frac{B\sigma - A\lambda}{\rho(\sigma^2 + \lambda^2)} + v_0 \right] \sin \lambda t + \left[ u_0 - \frac{B\lambda}{\rho} + \frac{B\sigma - A\lambda}{\rho(\sigma^2 + \lambda^2)} \right] \cos \lambda t \quad (14)$$

$$+ \frac{B}{\rho \lambda} - \left( \frac{\lambda^2}{\sigma^2 + \lambda^2} \right) \left( u_g + \frac{B}{\rho \lambda} \right) - \frac{\sigma \lambda}{\sigma^2 + \lambda^2} v_g$$

$$v = \left[ \frac{B}{\rho \lambda} - \frac{B\lambda - A\sigma}{\rho(\sigma^2 + \lambda^2)} - u_0 \right] \sin \lambda t + \left[ v_0 + \frac{B\sigma - A\lambda}{\rho(\sigma^2 + \lambda^2)} \right] \cos \lambda t$$

$$- \left( \frac{\sigma \lambda}{\sigma^2 + \lambda^2} \right) \left( u_g + \frac{B}{\rho \lambda} \right) + \frac{\lambda^2}{\sigma^2 + \lambda^2} v_g \quad (15)$$



माना  $f(x) = \frac{1}{x^2}$  है तो  $f'(x)$  का मान ज्ञात करें।

समाधान:  $f(x) = \frac{1}{x^2}$  है तो  $f'(x) = \frac{d}{dx} \left( \frac{1}{x^2} \right)$

$= \frac{d}{dx} (x^{-2})$

$= -2x^{-3}$

$\therefore f'(x) = -\frac{2}{x^3}$

$$\therefore \frac{d}{dx} \left( \frac{1}{x^2} \right) = -\frac{2}{x^3}$$

$$\therefore f'(x) = -\frac{2}{x^3}$$

उदाहरण 2: माना  $f(x) = \frac{1}{x^2}$  है तो  $f'(x)$  का मान ज्ञात करें।

समाधान:  $f(x) = \frac{1}{x^2}$  है तो  $f'(x) = \frac{d}{dx} \left( \frac{1}{x^2} \right)$

$$= \frac{d}{dx} (x^{-2})$$

$$= -2x^{-3}$$

$\therefore f'(x) = -\frac{2}{x^3}$

$$\therefore \frac{d}{dx} \left( \frac{1}{x^2} \right) = -\frac{2}{x^3}$$

उदाहरण 3:

$$\frac{d}{dx} \left( \frac{1}{x^2} \right) = -\frac{2}{x^3}$$

$$\therefore \frac{d}{dx} \left( \frac{1}{x^2} \right) = -\frac{2}{x^3}$$

उदाहरण 4:

$$\frac{d}{dx} \left( \frac{1}{x^2} \right) = -\frac{2}{x^3}$$



If  $A = 0$  and  $\gamma = \infty$  the pressure distribution becomes  $p = p_0 + By$ , which is the same as used in Chapter II of this paper and also in one of Forsdyke's [4] solutions for a particular time variation of pressure.

Substituting  $A = 0$  and  $\gamma = \infty$  in (14) and (15) we get

$$u = v_0 \sin \lambda t + (u_0 - \frac{B\lambda}{\rho}) \cos \lambda t - \frac{B}{\rho\lambda} \quad (16)$$

$$v = (\frac{B}{\rho\lambda} - u_0) \sin \lambda t + v_0 \cos \lambda t \quad (17)$$

Equations (16) and (17) are the same as Forsdyke obtained when the substitutions  $u^1 = u_0 - u_g$  and  $v^1 = v_0 - v_g$  are made since  $u_g = \frac{B\lambda}{\rho}$  and  $v_g = 0$  for  $A = 0$ ,  $\gamma = \infty$ .

Equations (16) and (17) differ from (3) and (4) in spite of the same pressure distribution. This is consistent since different restrictions were placed on the motion; namely, constant Coriolis parameter and independence of space for (16) and (17), while steady state, independence of longitude, and a variable Coriolis parameter in (3) and (4).

Numerical evaluation of equations (16) and (17) could be done by assigning values to the constants  $A$ ,  $B$ , and  $\gamma$  consistent with observed pressure patterns and then computing  $u$  and  $v$  as a function of latitudes and time. This evaluation was not undertaken here for lack of time.



## V. CONCLUSIONS

The equations of motion may be solved for equatorial regions only with rather restrictive assumptions even for relatively simple pressure fields.

Although the solutions in this paper give velocity fields resembling observed mean motions, and one commonly used restriction has been lifted, namely, the requirement non-divergence, they appear to be of little practical value in routine weather analysis. In addition to other limitations the necessary computations are laborious and time consuming.



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## CHAPTER IV

- 1. The first part of the chapter is devoted to a discussion of the various methods of determining the value of the constant  $k$  in the equation  $y = kx$ .
- 2. The second part of the chapter is devoted to a discussion of the various methods of determining the value of the constant  $k$  in the equation  $y = kx^2$ .
- 3. The third part of the chapter is devoted to a discussion of the various methods of determining the value of the constant  $k$  in the equation  $y = kx^3$ .
- 4. The fourth part of the chapter is devoted to a discussion of the various methods of determining the value of the constant  $k$  in the equation  $y = kx^4$ .
- 5. The fifth part of the chapter is devoted to a discussion of the various methods of determining the value of the constant  $k$  in the equation  $y = kx^5$ .
- 6. The sixth part of the chapter is devoted to a discussion of the various methods of determining the value of the constant  $k$  in the equation  $y = kx^6$ .
- 7. The seventh part of the chapter is devoted to a discussion of the various methods of determining the value of the constant  $k$  in the equation  $y = kx^7$ .
- 8. The eighth part of the chapter is devoted to a discussion of the various methods of determining the value of the constant  $k$  in the equation  $y = kx^8$ .
- 9. The ninth part of the chapter is devoted to a discussion of the various methods of determining the value of the constant  $k$  in the equation  $y = kx^9$ .
- 10. The tenth part of the chapter is devoted to a discussion of the various methods of determining the value of the constant  $k$  in the equation  $y = kx^{10}$ .











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